

Magnetic Resonance Sounding

A new technique for hydrogeophysics

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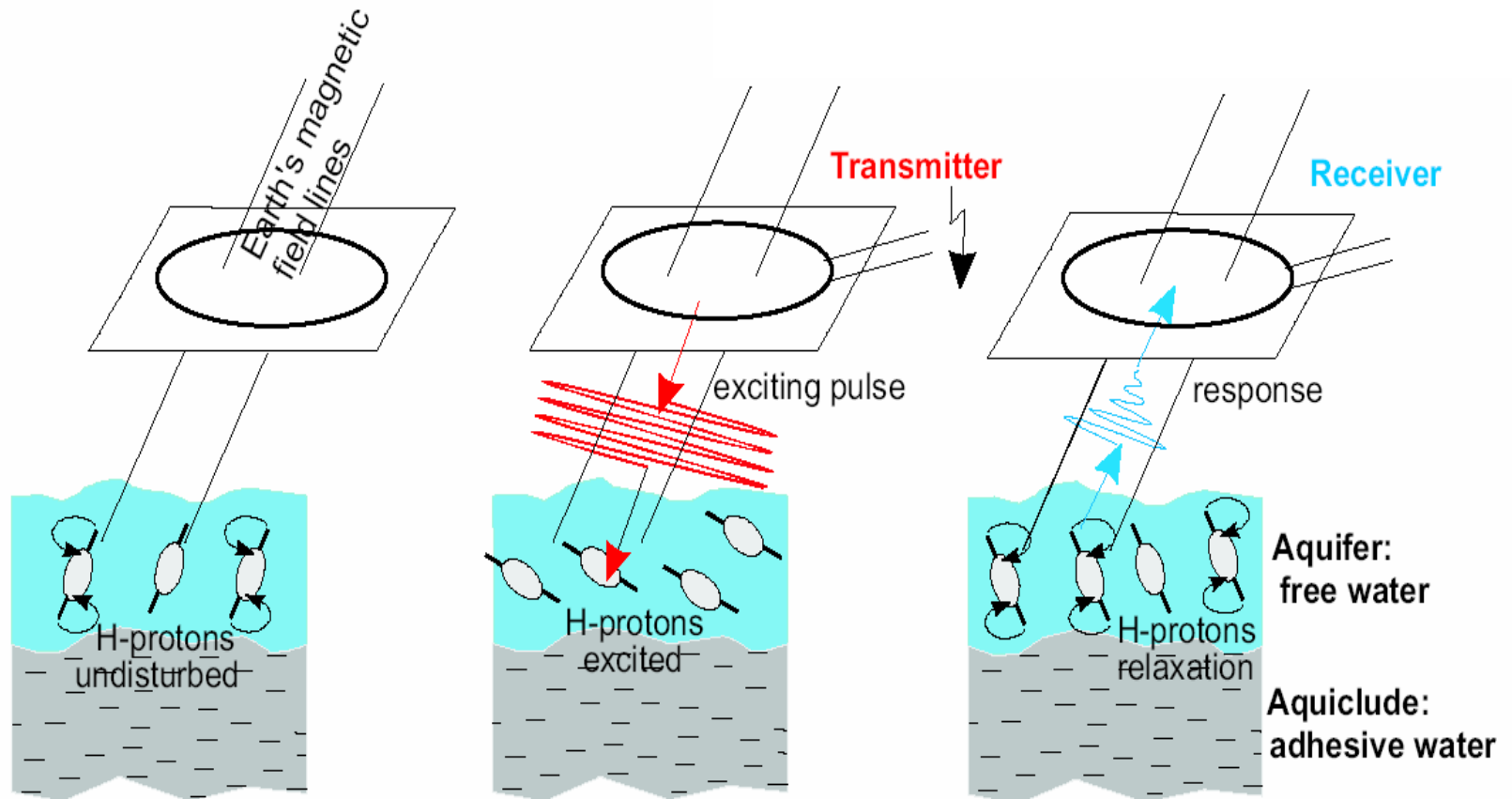
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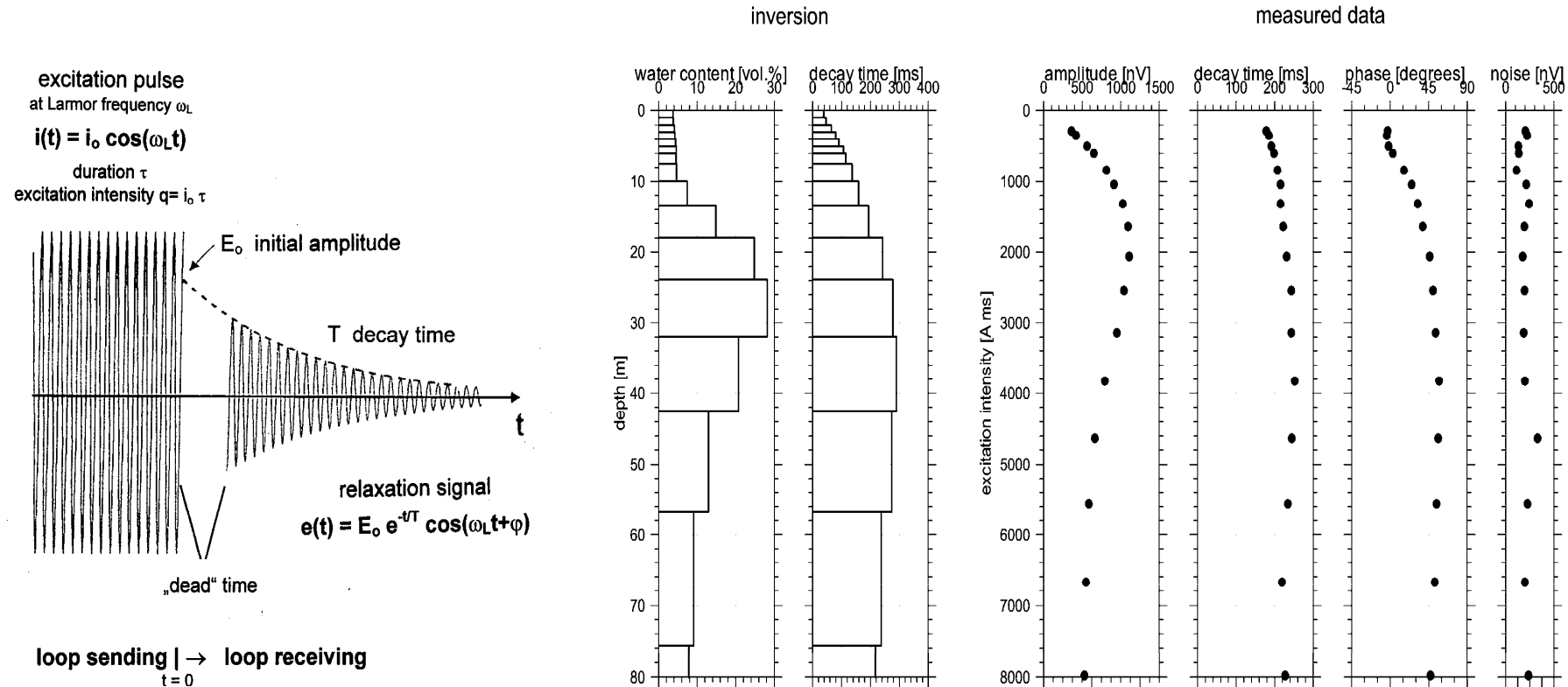
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Excitation and relaxation of hydrogen protons



Basic principle of MRS measurements



Measured quantity

E_0 Initial Amplitude

T Decay Time

φ Phase

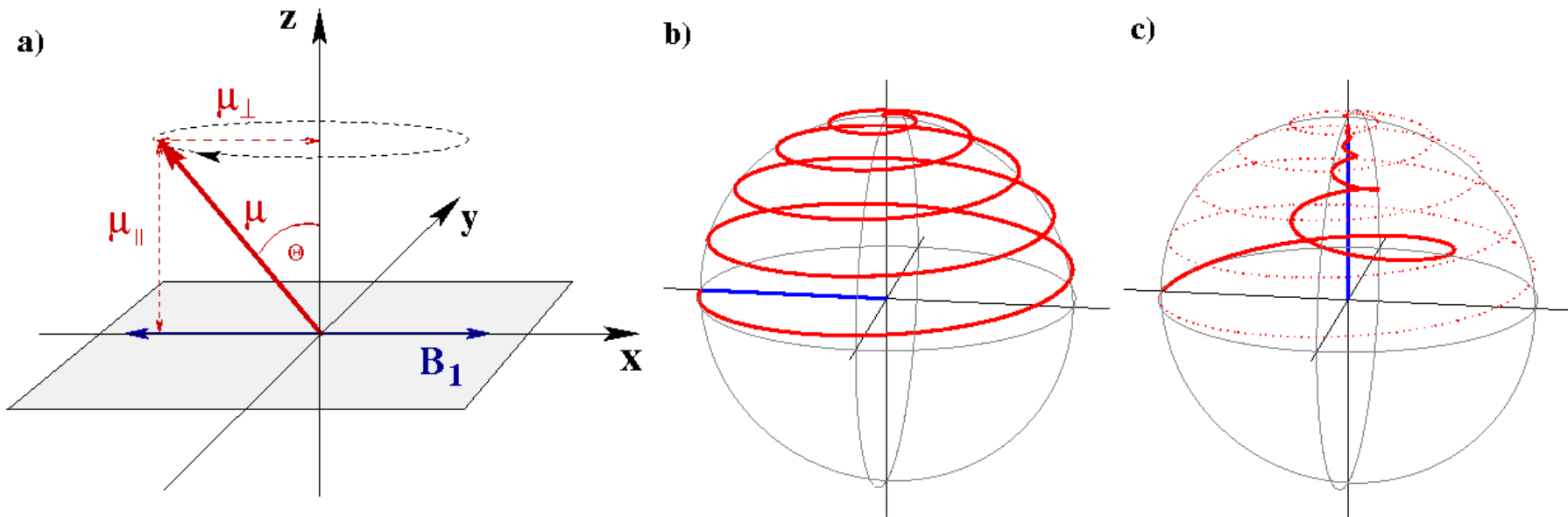
Derived parameter

\sim layer water content

\sim layer decay time \sim pore size \sim hydraulic permeability

\sim layer electrical conductivity

Spin behaviour during excitation and relaxation



(a) Components of the precessing magnetic moment of the Spin.

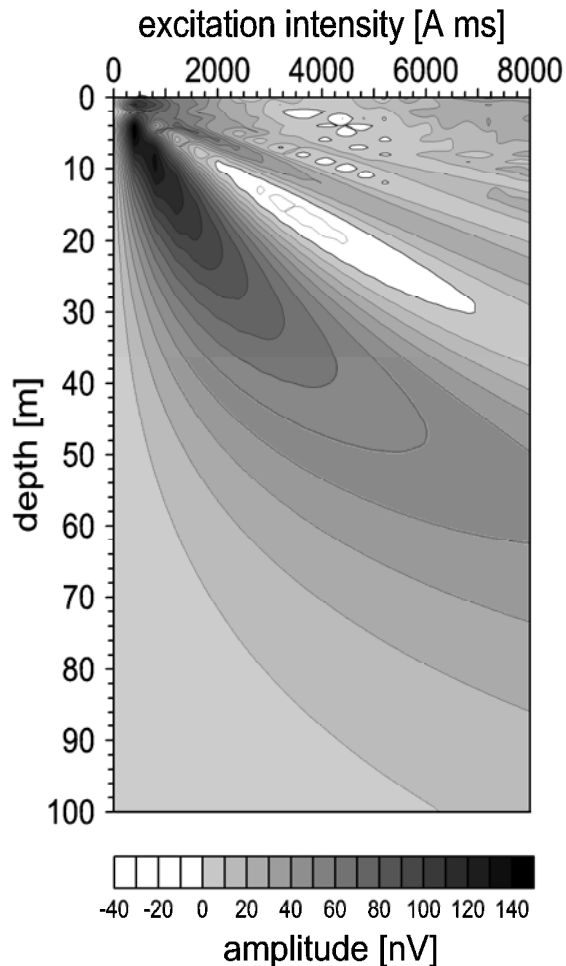
(b) Unit trajectory of the Spin magnetic moment during the excitation process

(c) and during the relaxation process

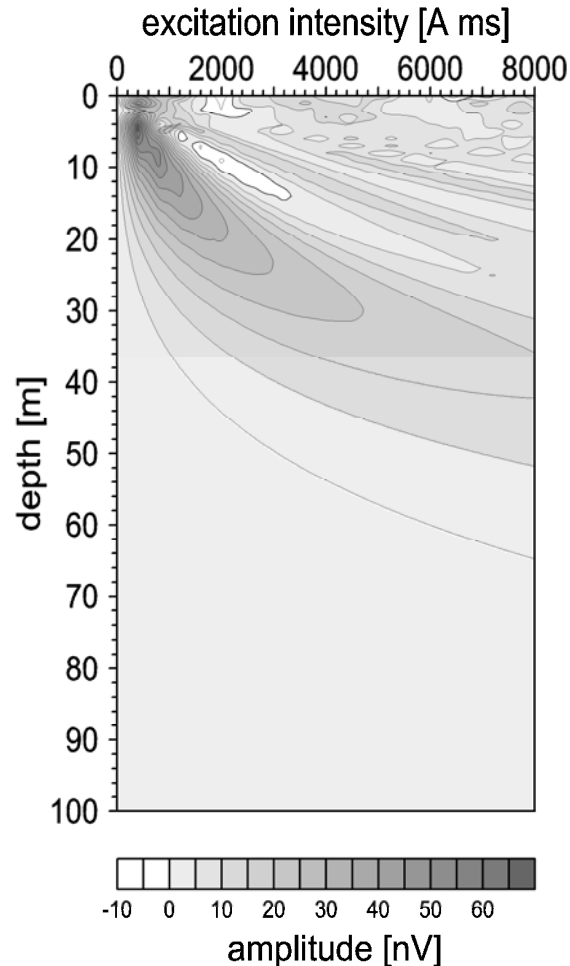
Direction of the static field B_0 is along the z-axis

Sensitivities at 1D soundings for different loop sizes

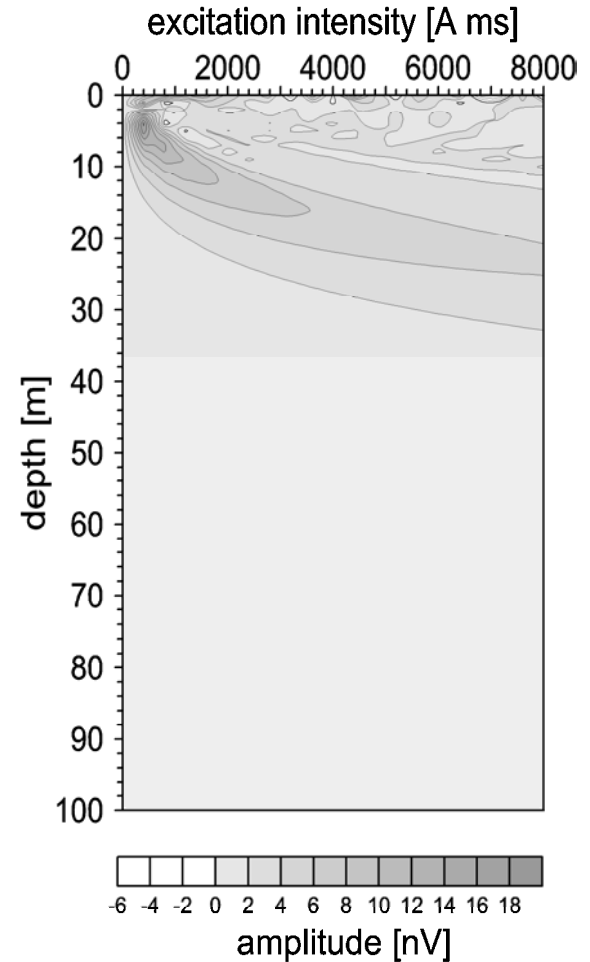
$d = 100$ m



$d = 50$ m



$d = 20$ m



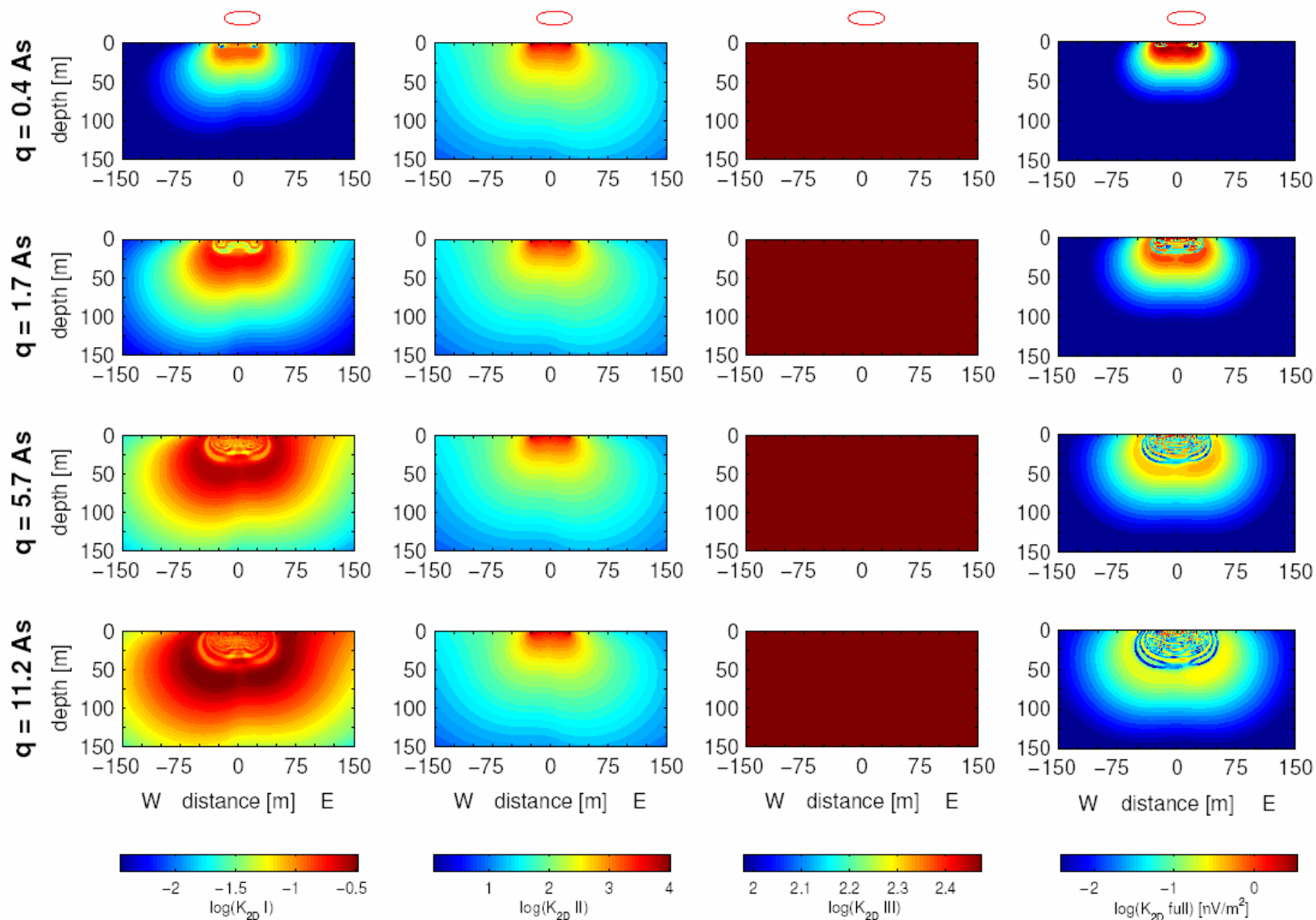
Basic equation for MRS response signal

Initial voltage in receiver loop

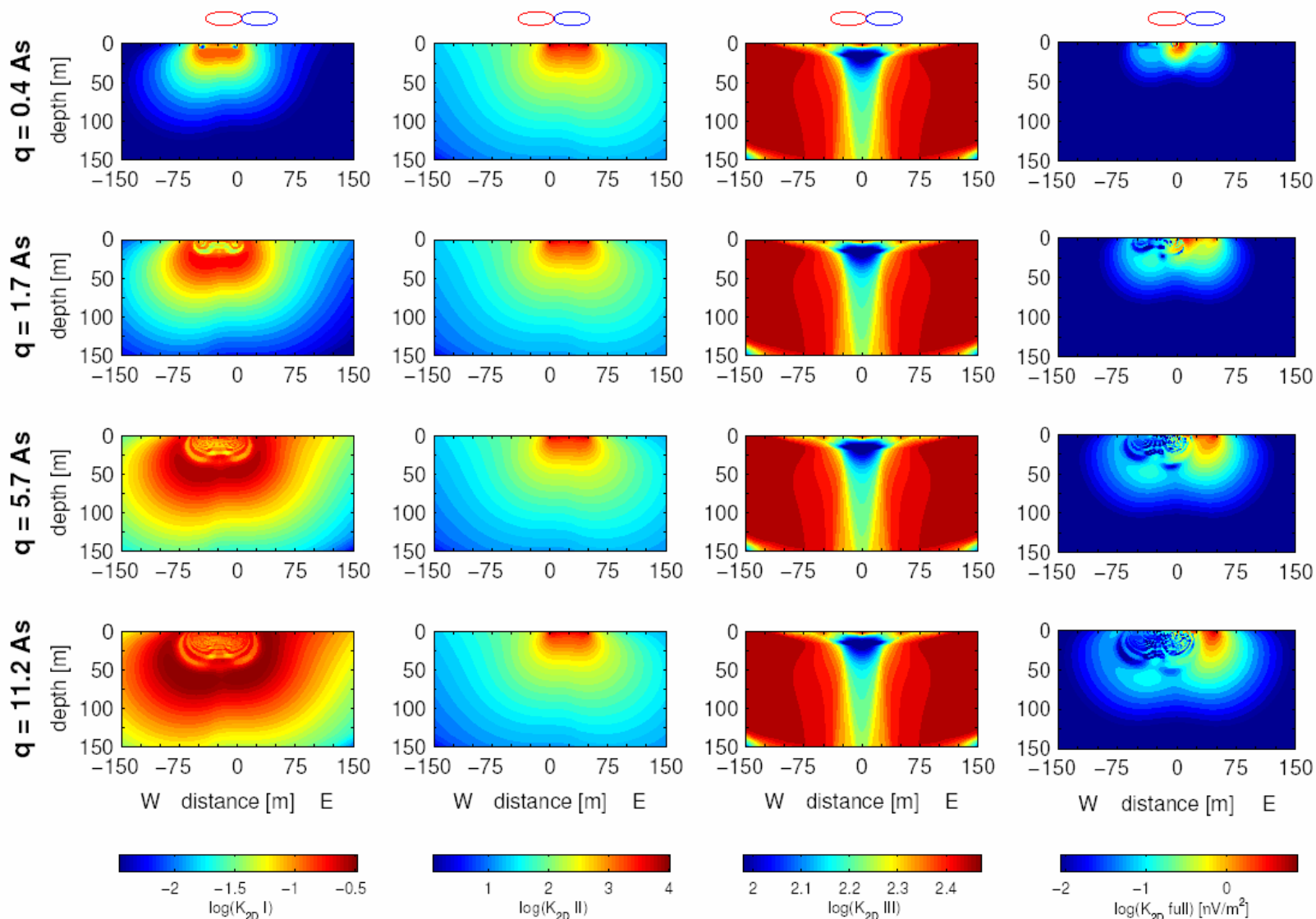
$$\begin{aligned}
 V_R(q) = & \omega_L \int d^3\mathbf{r} \left| \mathbf{M}_N^{(0)}(\mathbf{r}) \right| \sin(q\mathbf{B}_T^+(\mathbf{r})) \\
 & \times \mathbf{B}_R^-(\mathbf{r}) \cdot e^{i[\zeta_T(\mathbf{r}) + \zeta_R(\mathbf{r})]} \\
 & \times \left[\hat{\mathbf{b}}_R(\mathbf{r}) \cdot \hat{\mathbf{b}}_T(\mathbf{r}) + i\hat{\mathbf{B}}_0 \cdot \hat{\mathbf{b}}_R(\mathbf{r}) \times \hat{\mathbf{b}}_T(\mathbf{r}) \right]
 \end{aligned}$$

- (I) The Spin signal excitation induced by the transmitter loop
- (II) The sensitivity of the receiver loop to the excited Spin
- (III) The geometric relation of the magnetic fields of both loops

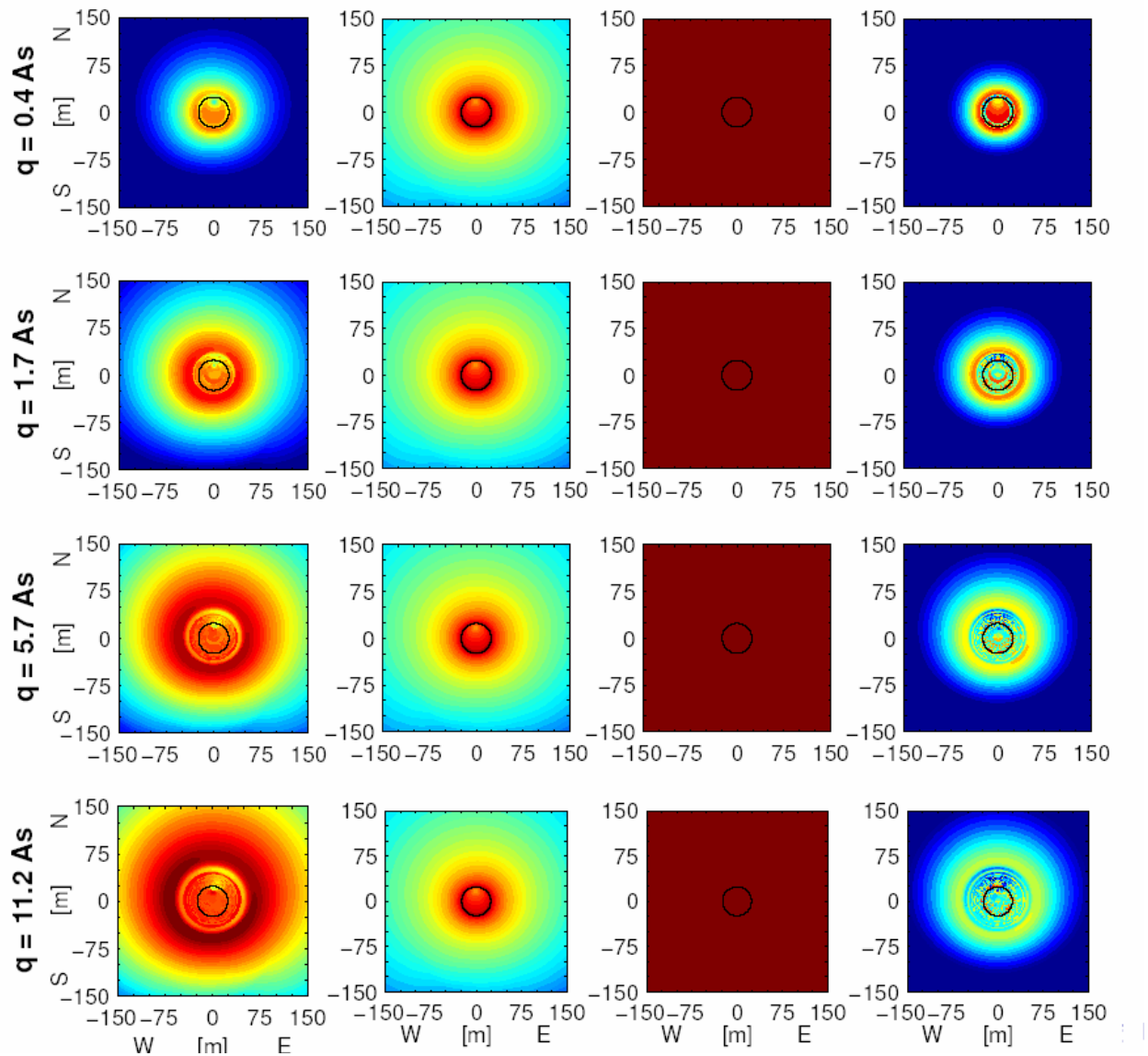
2d-sensitivities for coinciding receiver and transmitter



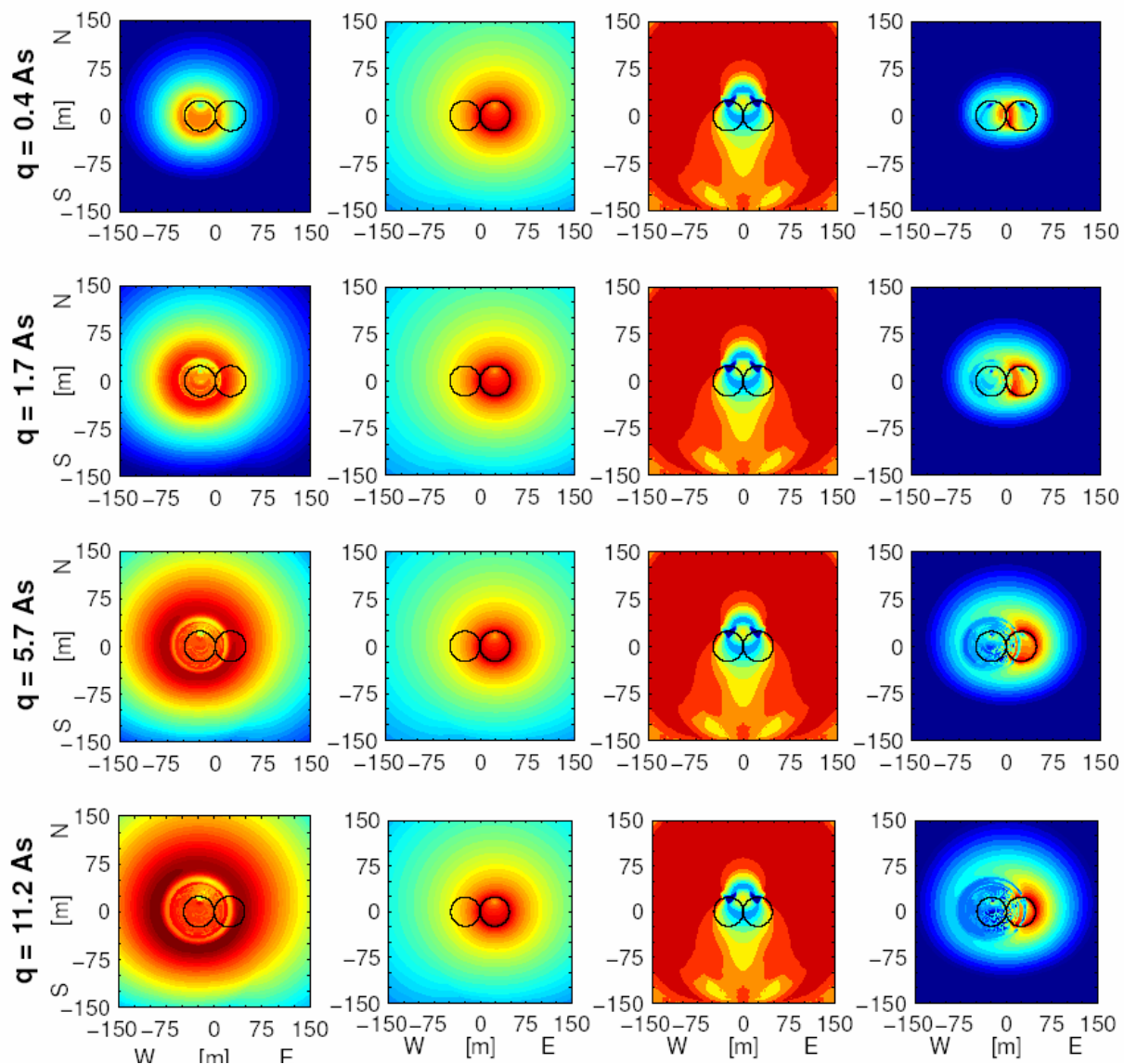
2d-sensitivities for seperated receiver and transmitter



2d-sensitivities for coincident receiver and transmitter

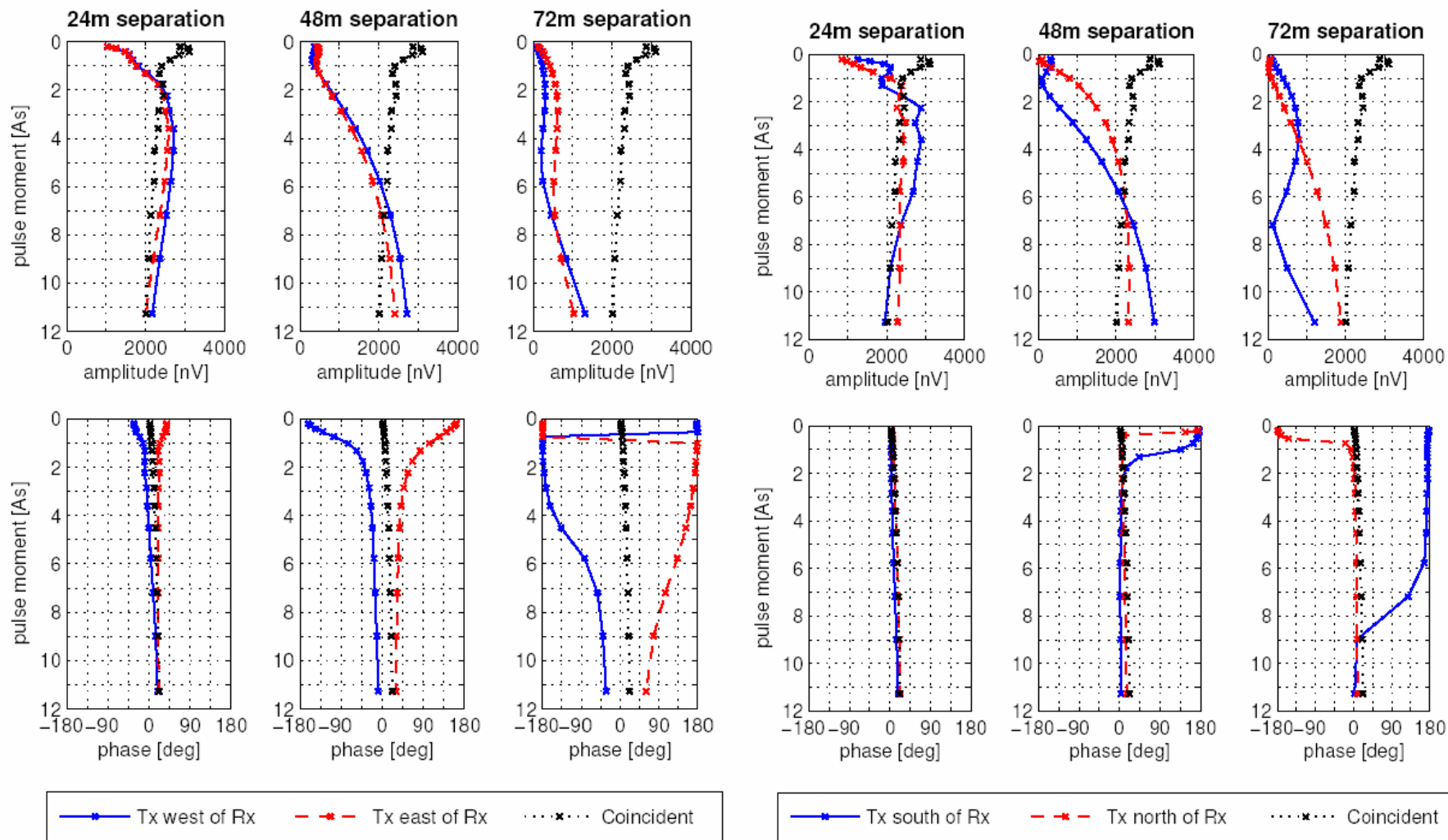


2d-sensitivities for seperate receiver and transmitter



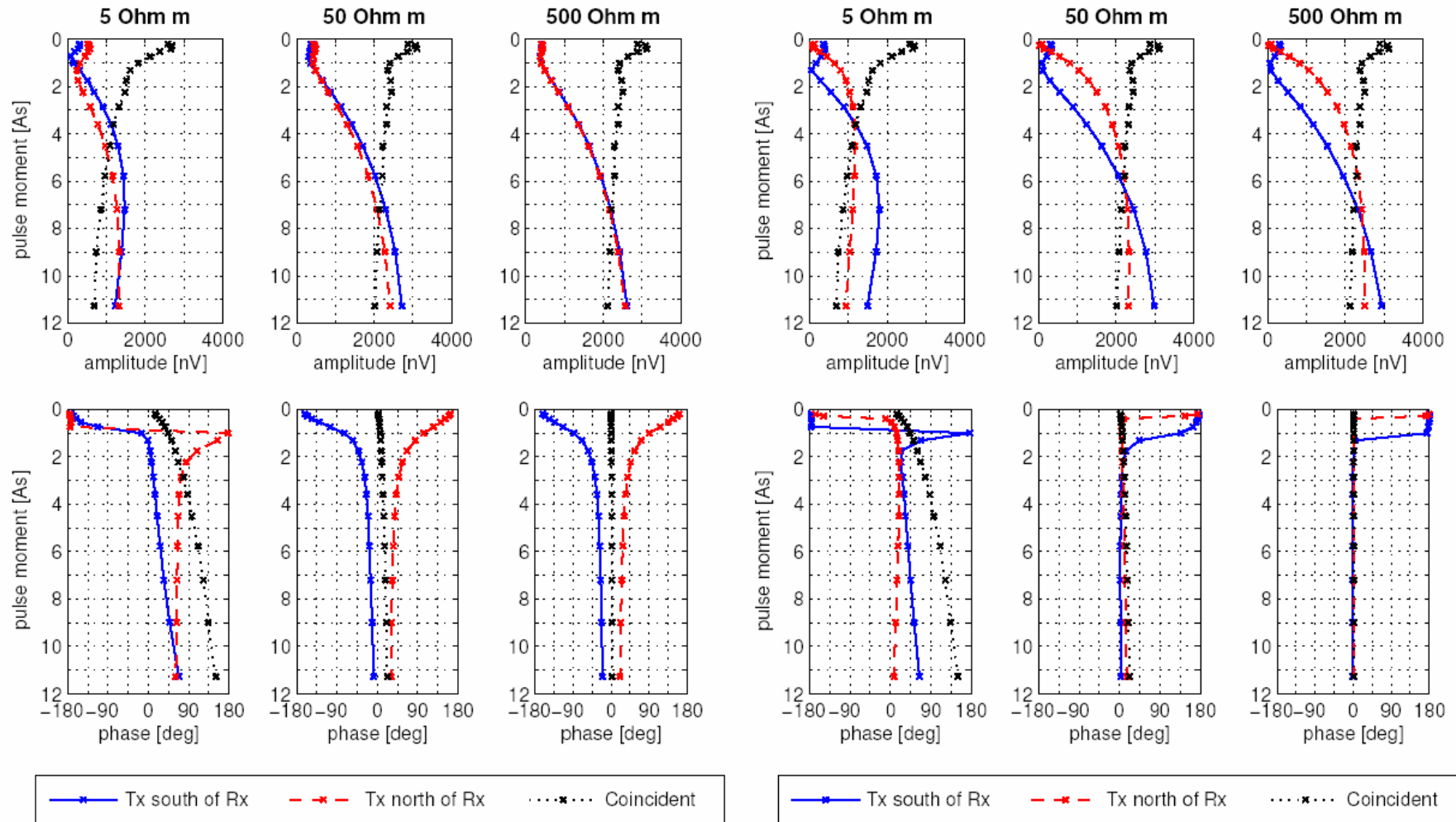
Synthetic data for homogeneous half space

effect of receiver and transmitter separation and separation direction

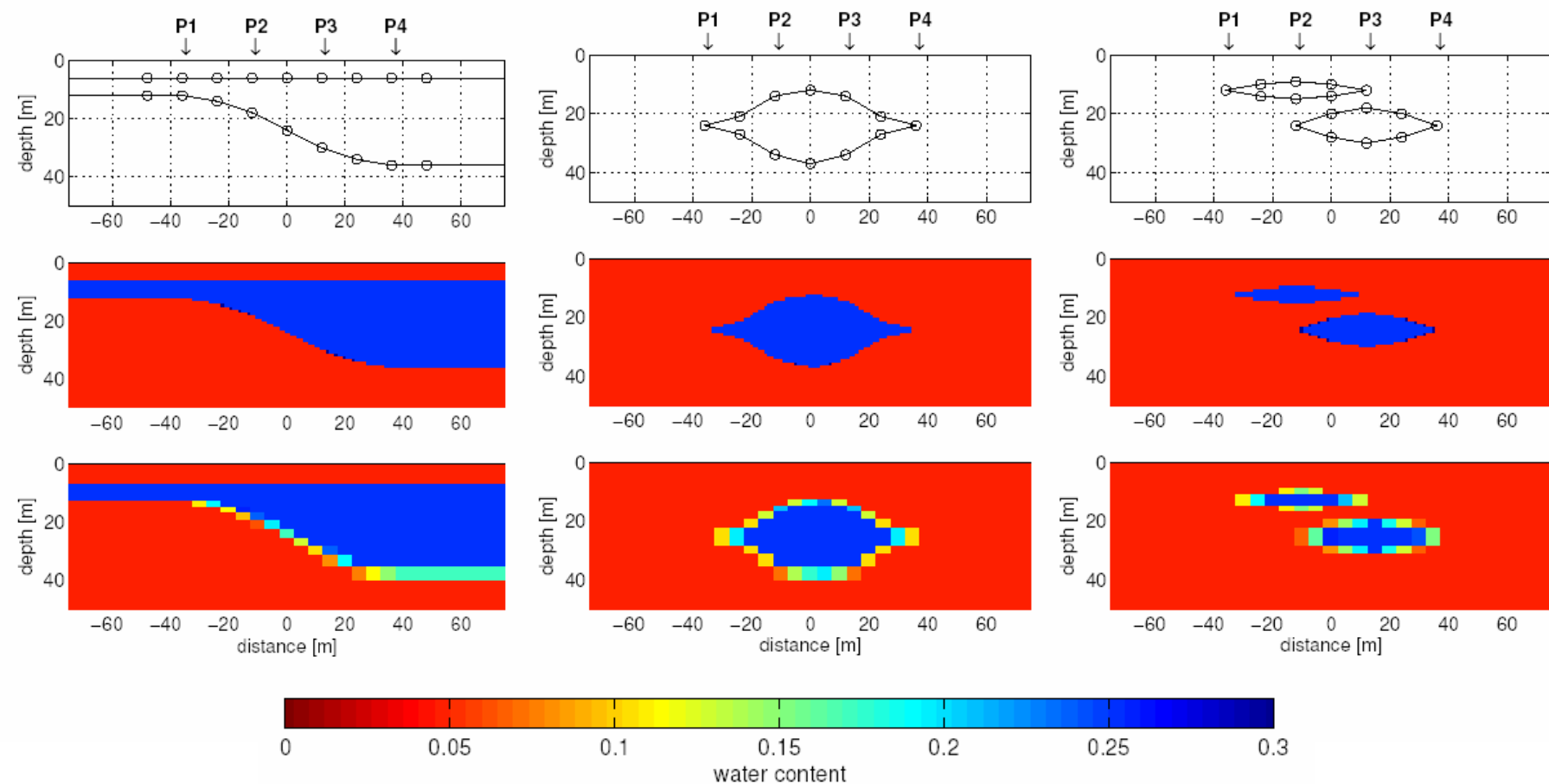


Synthetic data for homogeneous half space

effect of receiver and transmitter separation and resistivity



Models for tests with synthetic data

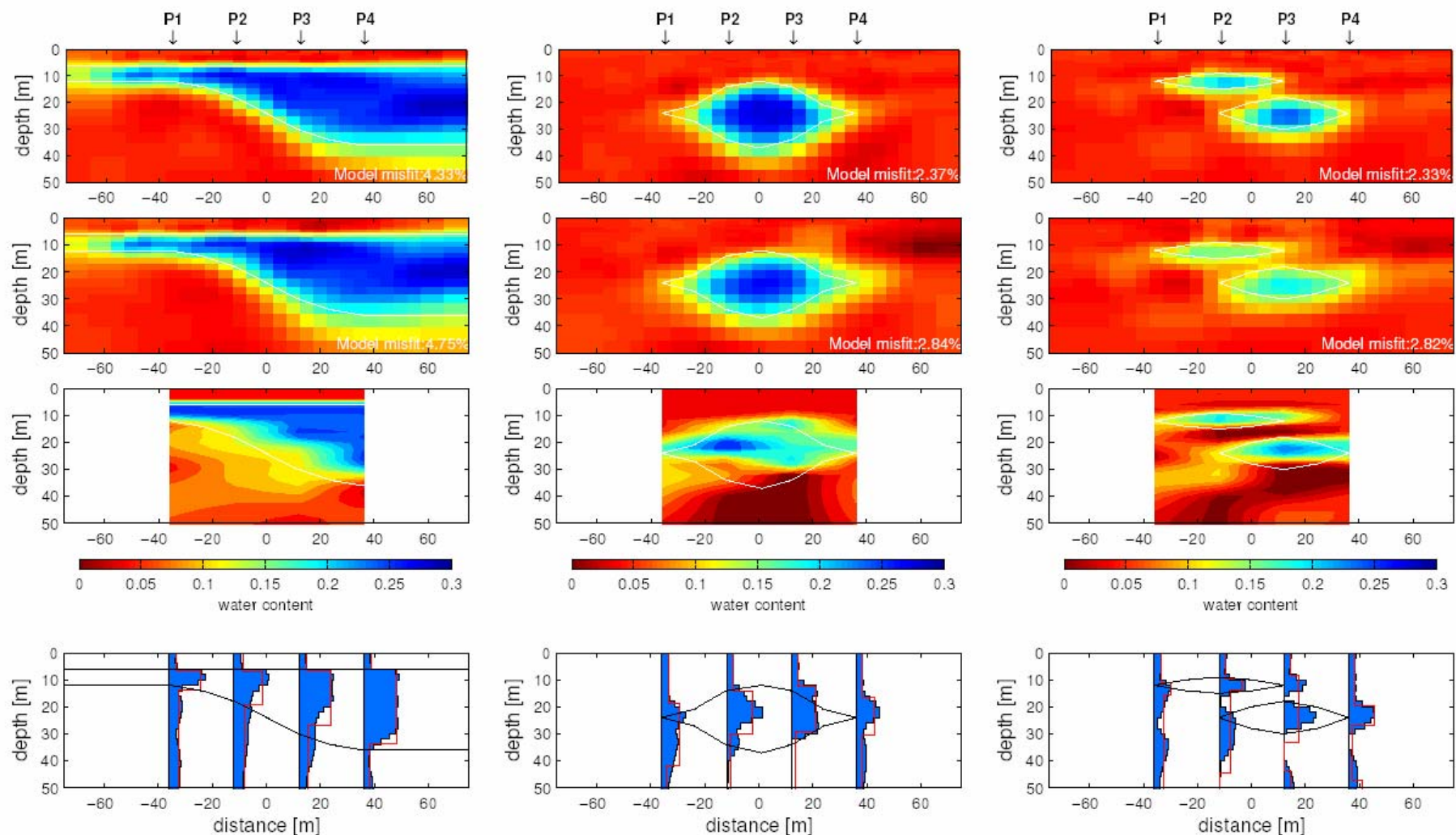


Results of inversion of synthetic data

(a) 2D full inversion with full set data of separated loops

(b) 2D full inversion with data of coincident loops

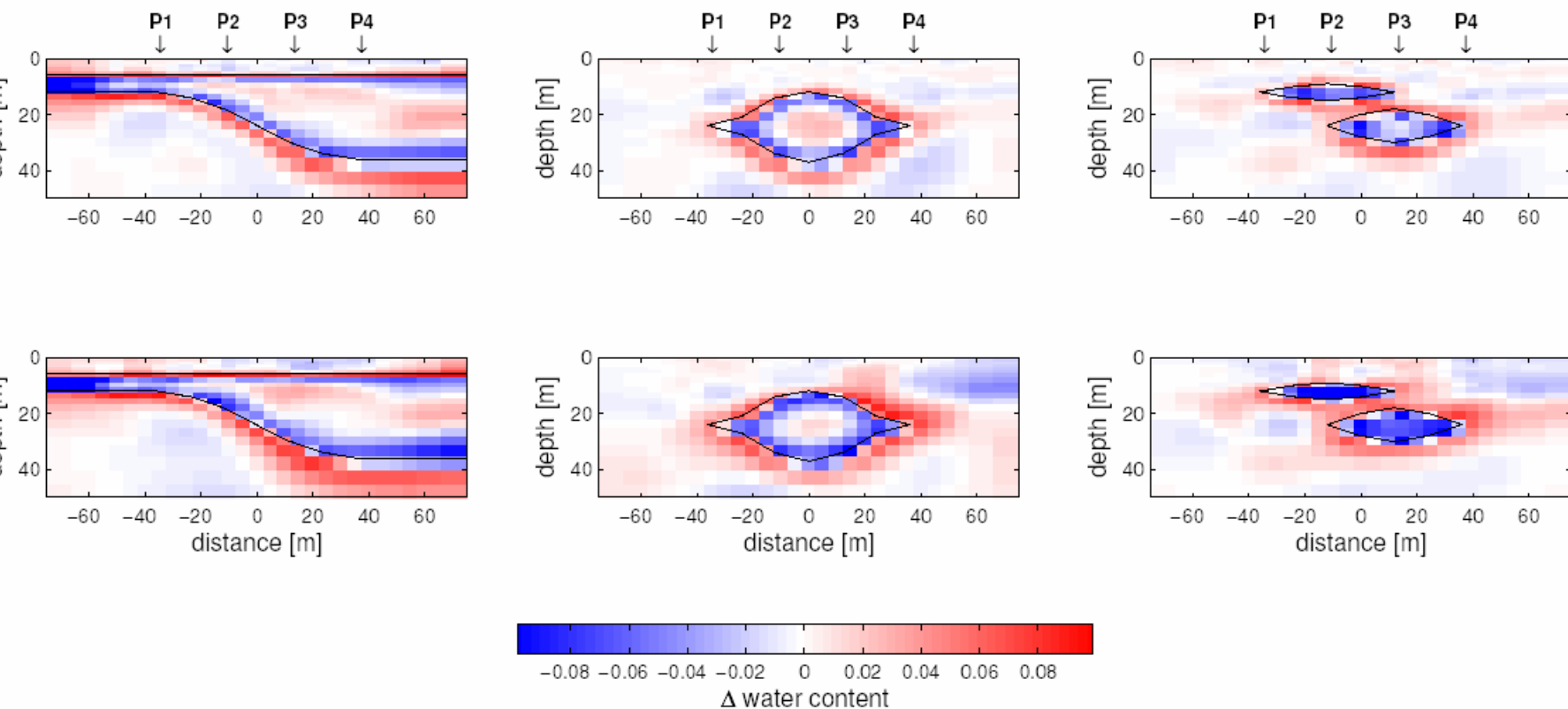
(c) 1D inversions in contourplot and (d) in usual sounding plots



Residuals of inversion of synthetic data

(a) 2D full inversion with full set data of seperated loops

(b) 2D full inversion with data of coincident loops

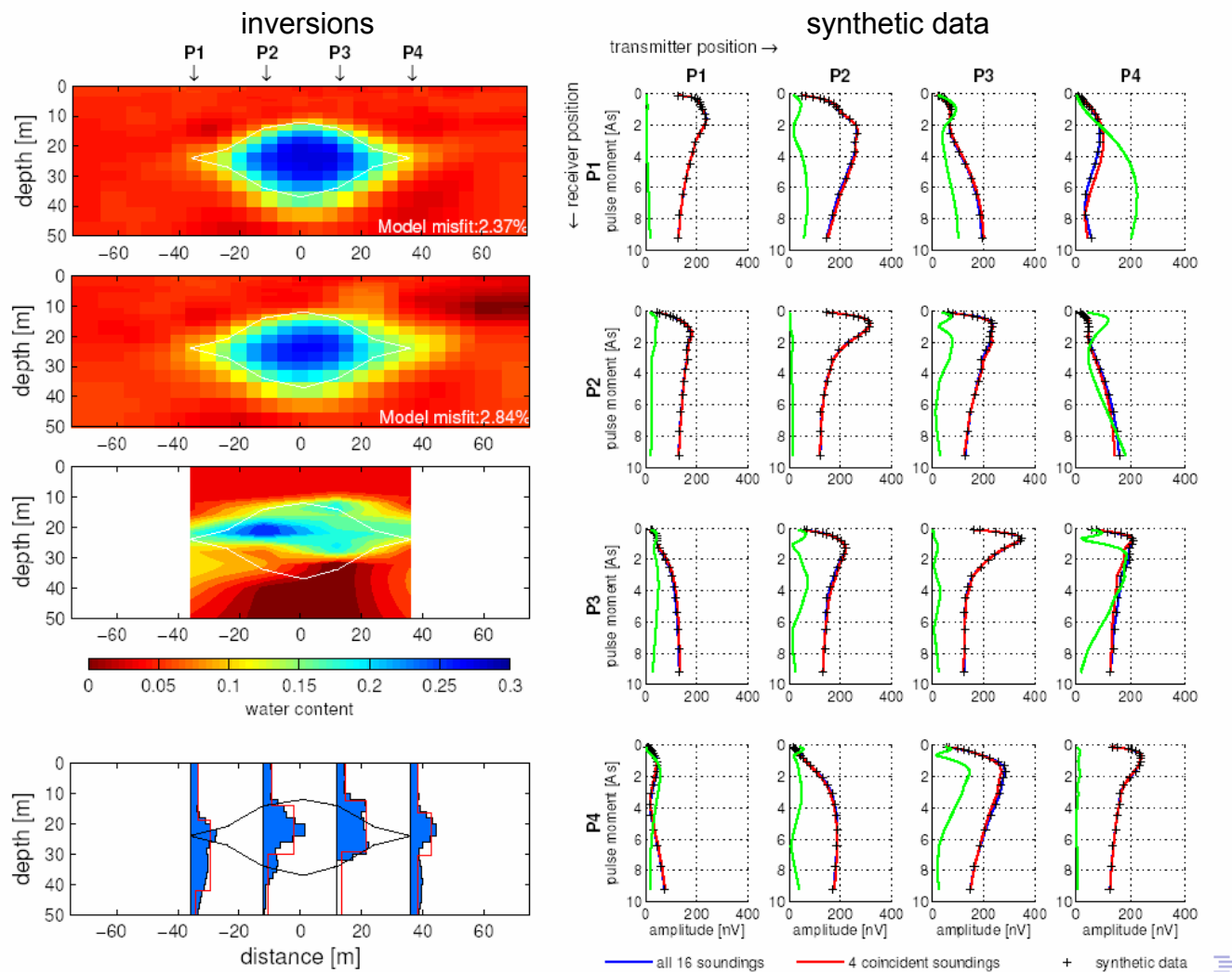


Synthetic data and inversion for a lens model

(a) 2D inversion with full set data of separated loops

(b) 2D inversion with data of coincident loops

(c) 1D inversions in contourplot and (d) in usual sounding plots

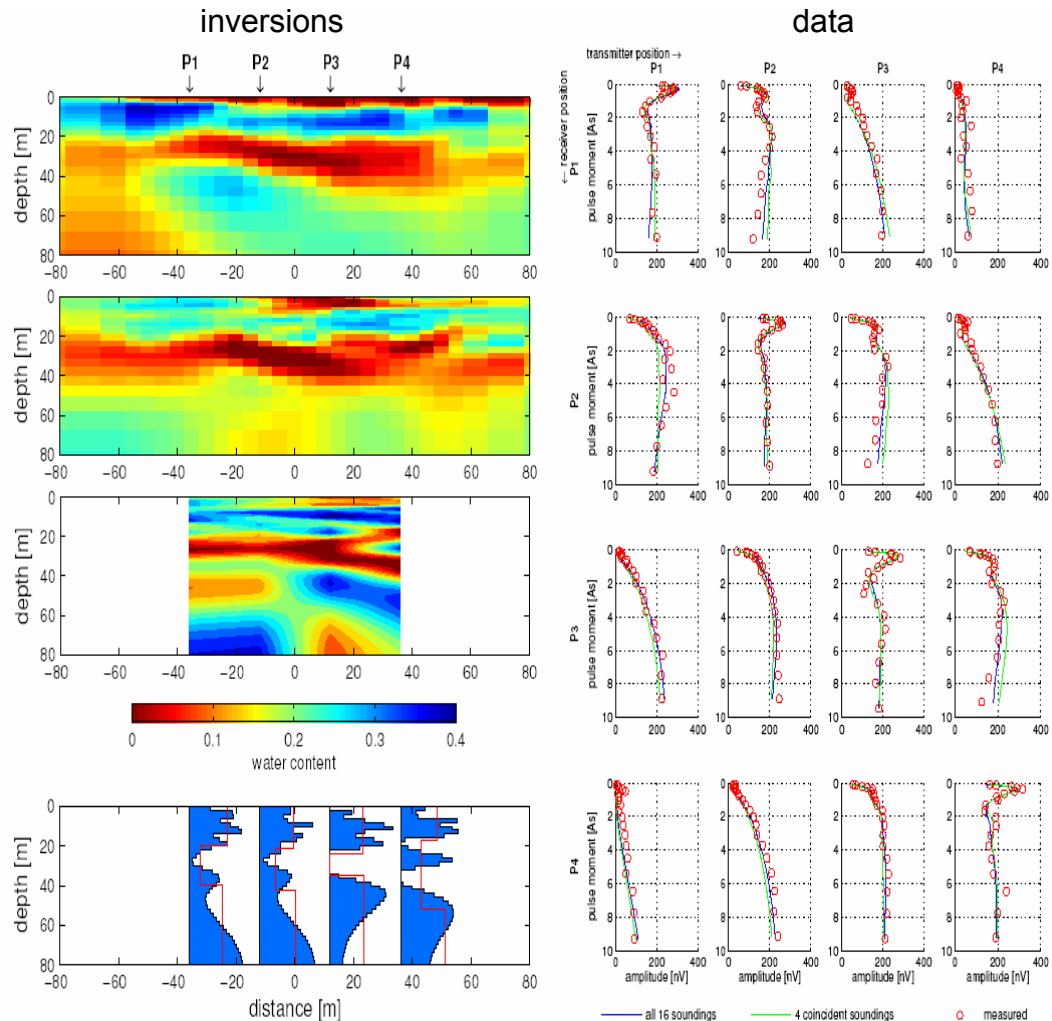


MRS measurements and inversion at test site 2

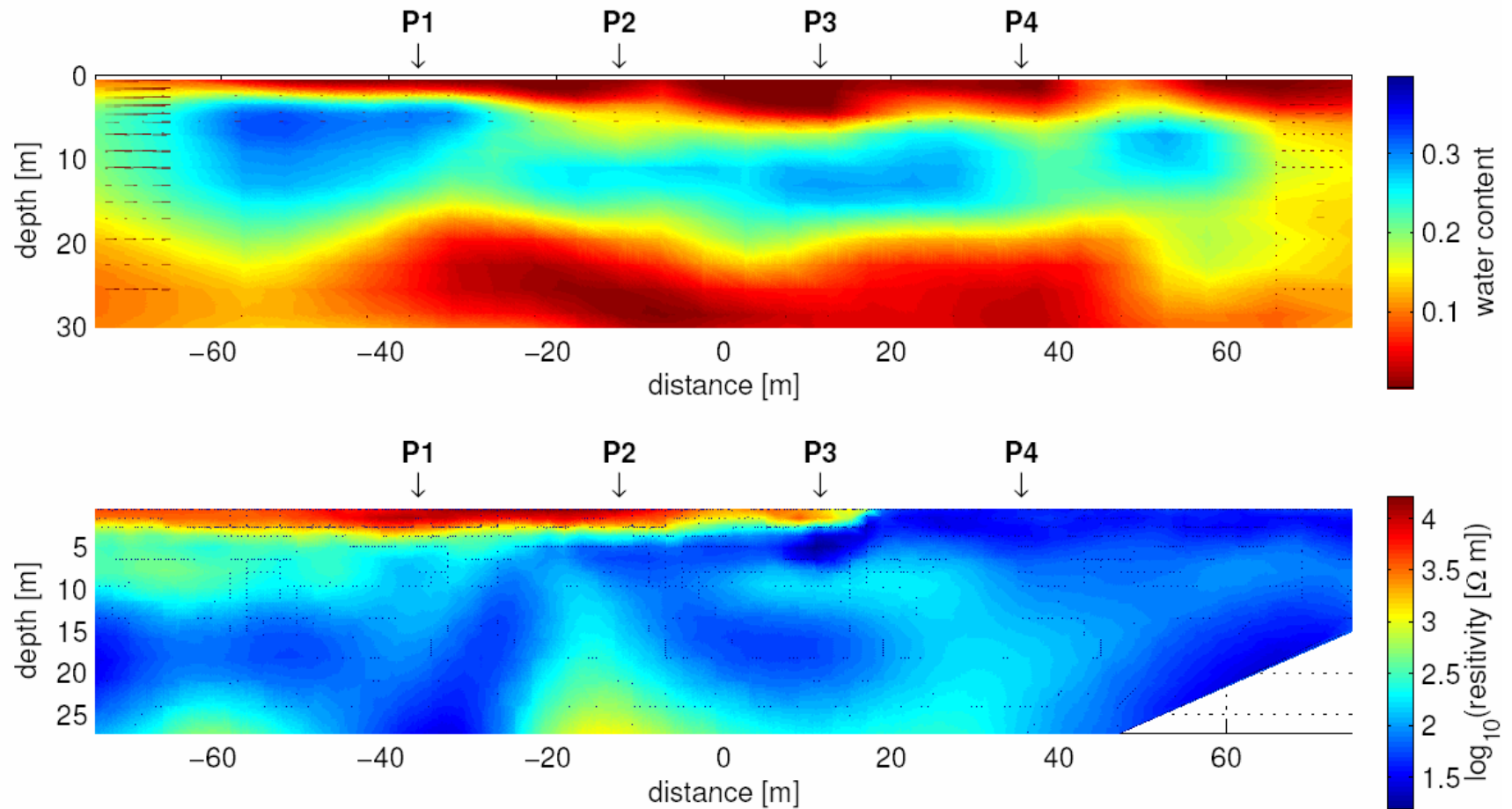
(a) 2D inversion with full set data of seperated loops

(b) 2D inversion with data of coincident loops

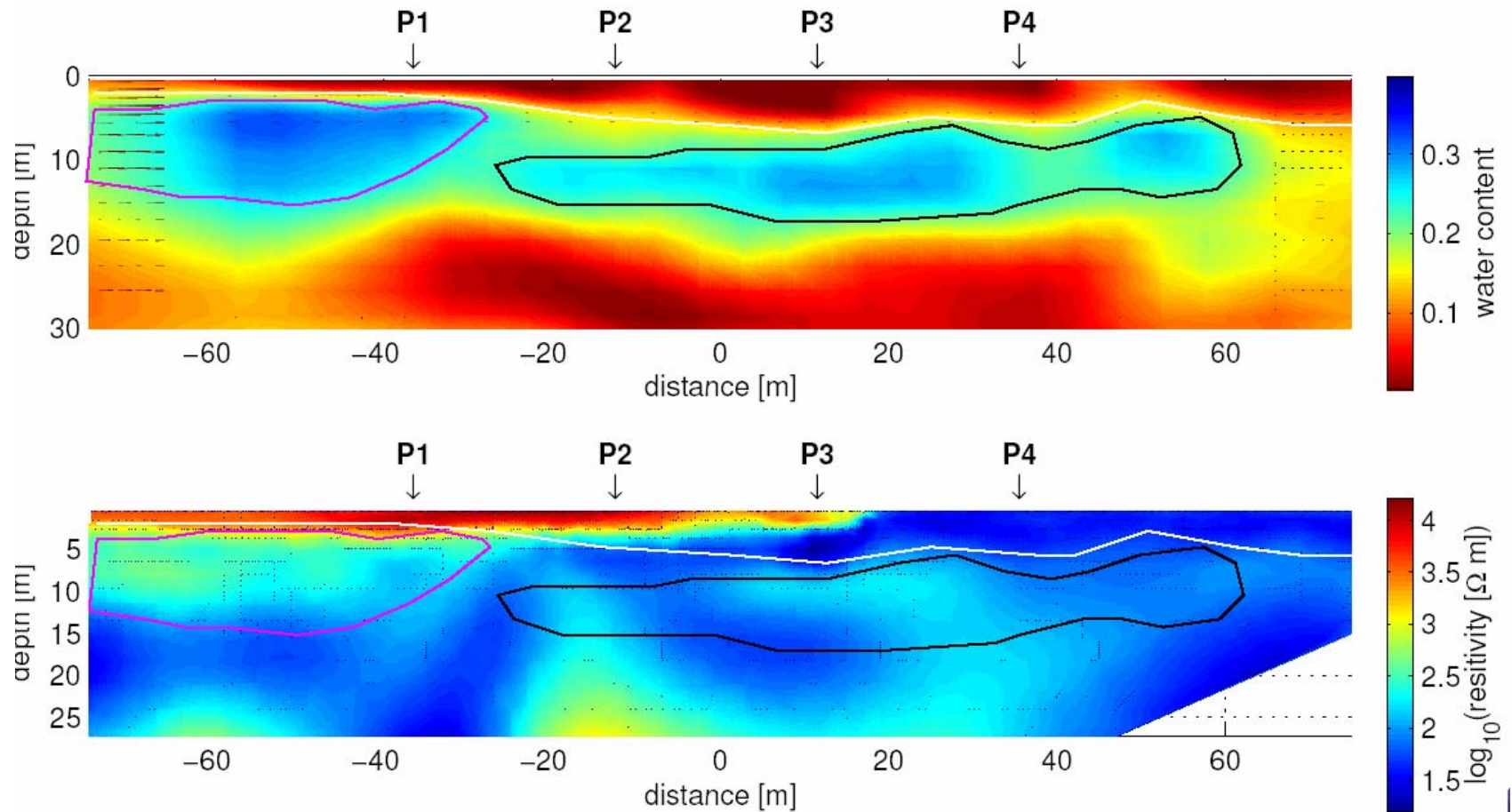
(c) 1D inversions in contourplot and (d) in usual sounding plots



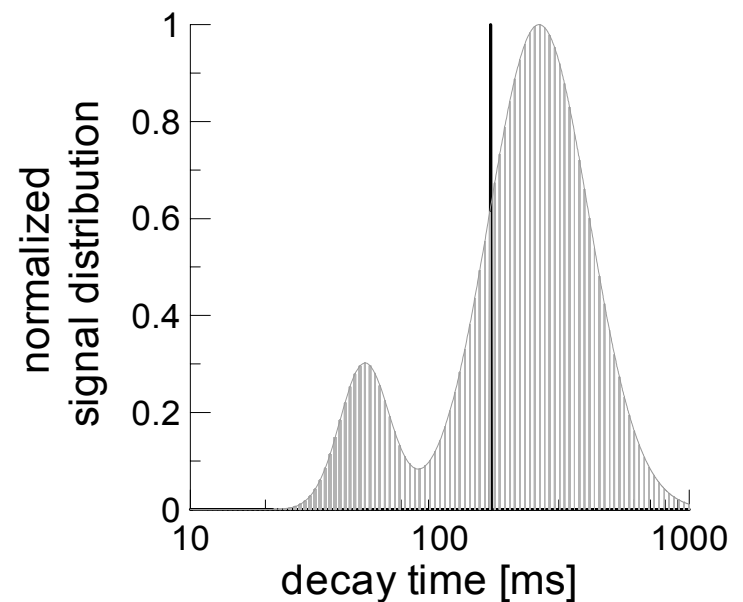
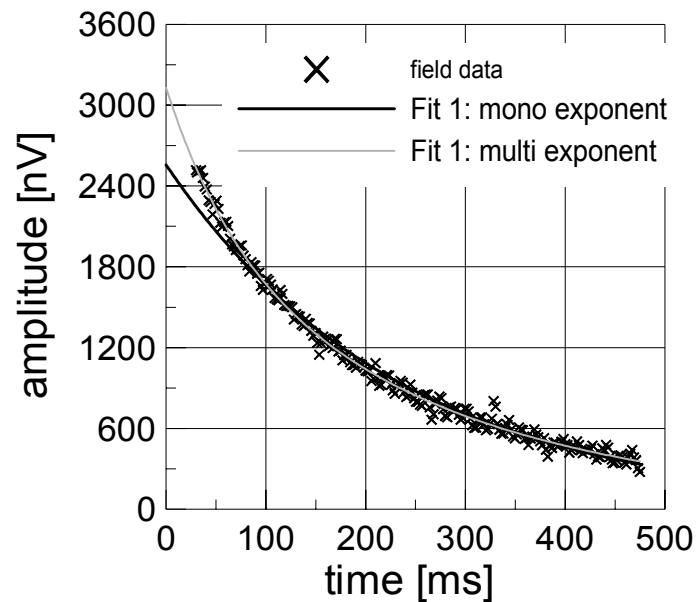
MRS and geoelectrics at test site 2



MRS and geoelectrics at test site 2

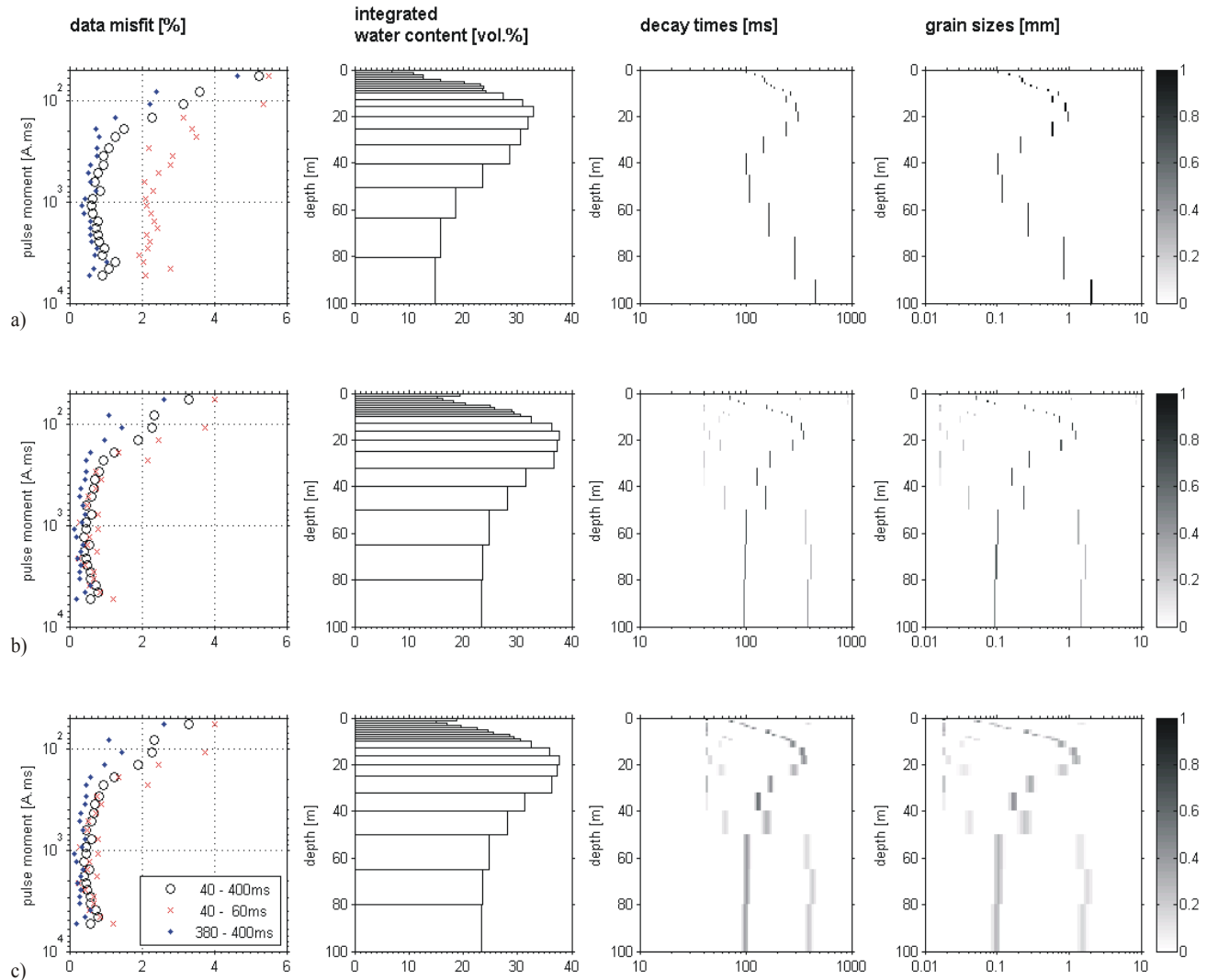


Decay times and distributions in MRS



Analysis of MRS decay times

(a) Monoexponential, (b) Biexponential, (c) Multiexponential



In conclusion

Magnetic Resonance Sounding (or – Tomography) (MRS and MRT) have passed the experimental stage and turned out to be a very useful method in particular for hydrogeophysics

Water content distribution can be detected explicitly and quantitatively

Effective pore sizes or even pore size distributions and followingly hydraulic permeabilities can be estimated

The potential of MRS can be increased in joint use, interpretation and inversion with other geophysical methods in particular with those of geoelectrics and electromagnetics